

# Hedrogeological Studies on Wadi El-Farigh Area, West Nile Delta, Egypt

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## ABSTRACT

The present study deals will deal with the evaluation of groundwater potentiality by using Hydrogeological and Hydrochemical investigation tools. The area of study is located south of Wadi El-Natron depression, southwest of the Nile Delta and left the Cairo – Alexandria Desert Road. It extends in a WNW-ESE direction for about 90 km with an average width of about 10 km. As calculated by the present author Wadi El-Farigh depression has a catchment area of 1112.5 km<sup>2</sup> enclosed between the contours of 120 m above sea level in the south and west and 4 m below sea level in north center part. It lies between latitudes 30° 00' 00" and 30° 30' 00" N and Longitudes 30° 00' 00" to 30° 50' 00", E. The area includes lands belonging to the two governorates of Behira and Giza. The low area of wadi El-Farigh depression is dominated by sand accumulation and rock fragment. The aquifer system is belonging to Moghra (Lower Miocene) and the analyzed water samples of the investigated area show fresh to brackish water. The T.D.S. for the analyzed water samples ranges from 252 ppm to 2572 ppm. The depth to water in Miocene aquifer varies from 49 m at the area close to El Rayah El Naseri to 176.88 m at high topographic areas towards the west. The total salinity of the Miocene aquifer (Moghra aquifer) varies from fresh to slightly brackish water.

**Keywords:** Groundwater-Aquifer- Nile Delta.

## I. INTRODUCTION

The areas under investigation occupy a portion of the Western Nile Delta region (Fig. 1). It lies between latitudes 30° 00' 00" and 30° 30' 00" N and Longitudes 30° 00' 00" to 30° 50' 00", E. It is considered as a depression of low relief (El-Tahrir gravel plan, Wadi El-Natron depression and Wadi El-Farigh Depression).

The available studies are dealt with the geomorphology, geology, hydrogeology, hydrology and hydrogeochemistry. Shata (1947), (1953), and (1962), El Fayoumy (1967), Attia (1975) (1982), Picard (1955), Said (1962), Seanad (1973) & Omara and Sanad (1975), El Shazly et al., (1975), General Petroleum Company "GPC" (1977, El Ghazawi and Atwa, 1994), El-Abd (2005), Shafeek (2015).



**Figure 1.** Location of the study area

The study of climatic conditions is of paramount importance for the study of geomorphological elements due to the impact of climate on the composition of the external factors. The current study is based on weather

station wadi El-Natron as the nearest weather station for the wadi El-Farigh depression.

- The average annual temperature of 21.5 ° C and the average ranges between 17.3 ° C for the winter months and 25 ° C for the summer months, suggesting that depression moderate climate in winter and hot in summer.
- Average annual wind speed (17.7 km / h), and this rate ranges between (16.3 km / h), and winter (19.1 km / h) for the summer and recorded the highest speed of the wind during the year in the month of April (21.1 km / h), while the lowest speed of the wind in the month of January and December (15 km / h)
- The annual average relative humidity in the wadi El-Farigh depression 55%, ranging from 58% for the winter months and 52% for the summer months.
- The monthly rate of evaporation ranges between 5.2 mm / day (average month of January), and 14.1 mm /day (average month of June), and up the annual rate of evaporation of 9.4 mm day.
- The annual average for the Rainfall that falls on the depression 3.5 mm. this quantity is of great significance to the emergence of depression and wadi El- Farigh, and suggest that the depression has arisen under moist conditions in ancient times was more rainfall than it is now, and maybe it happened in the Middle Miocene.
- These aridity index results indicate typical desert conditions for the study area. Form all the main Hydrometeorology finding we can conclude that the study area is characterized by a hyper- arid climate.

## II. Regional Geomorphology

The area under consideration is of Depression and low relief (El-Tahrir gravel plan, Wadi El- Natrun depression and Wadi El-Farigh Depression). The general land surface slopes gently to the northern and eastern directions. Landforms are classified into 3 geomorphic units (Fig. 1.2). These geomorphic units from north to south are:

**The alluvial plains:** The alluvial plains constitute one of the most striking land features in the study area. These plains extend between Rosetta branch and the eastern fringes of Maryut tableland. They are classified into young and old alluvial plains.

**The structural plains:** The structural plains occupy a wide area to the south and west of the old alluvial plain. It consists of a number of alternating structural ridges and structural depressions, reflecting the impacts of both the Lithologic and structural factors. Old gravel surface stretch for a long distance on the gentle slopes of the structural plain as well as the escarpment bounding Wadi El Farigh on the northern side. They are classified into structural ridges and structural depressions.

**The shifting sand:** Drift sands and sand sheets cover wide portions of the old alluvial plains and the lowest parts of the large depressions. Also there are a series of long, narrow parallel elongate sand dunes known as El Heneishat sand dune chains. These chains extend southwards till the northern periphery of El Faiyum depression.

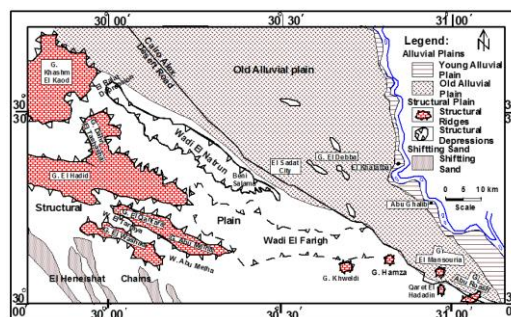


Figure 2. Geomorphologic map of west Nile Delta, Egypt, (Compiled after different authors).

## III. Stratigraphy

The study area is covered by extensive sedimentary exposures ranging in age from Late Cretaceous to Quaternary. In subsurface the sedimentary section has a thickness of about 4000 m resting on the basement rocks. The Late Cretaceous sediments have a localized occurrence on the crest of the complicated folded structure to the west of El Giza (Abu Rawoash anticline). Eocene and Oligocene sediments are of limited distribution in the environs of Cairo (Fig. 3) and (Fig.4).

**Miocene sediment:** Miocene sediment is represented by El Moghra Formation, which covers the south and west of Wadi El Natrun and the area from Wadi El Farigh in the east to El Qattara depression in the west. In Wadi El Farigh the thickness of Miocene sediments is about 150 m, which is rapidly increases to about 900 m in the northwest direction, and the facies of Miocene become more marine in that direction. Moghra Formation is mainly composed of coarse sands, sandstone and clay

interbeds with vertebrate remains and silicified wood, which becomes gravelly at base. The concerned formation rests unconformable on basalt sheets in several localities at surface as well as in the subsurface. Basalt sheets are located at levels ranges from -767 m below mean sea level at the southwest of Wadi El Natrun to 158 above mean sea level at Qaret El Haddadin. It located at depths ranges from ground surface to 839 m. They are located at different levels due to the effect of many faults in the form of horst-like and graben-like structures. Miocene lithofacies are Sandy gravel, gravelly sand, slightly gravelly sand and sand. Environments of deposition are different all over the Miocene aquifer from turbidity currents, fluvial and beach. The fluvial environments are the dominant environments.

The heavy minerals assemblage of El Moghra sediments have been recycled from older sediments which supplies the highly stable detritus (ZTR) and high metamorphic and acid igneous rocks of the basement rocks of the Eastern Desert mountains through the Eonile river tributaries which drained the Red Sea mountains during the Miocene period

**Pliocene sediments:** Pliocene sediments have a wide distribution in Wadi El Natrun depression and its vicinities; it is divided from base to top into Wadi El Natrun Formation and El Hagif Formation. Wadi El Natrun Formation is classified into Muluk and El Solimaniya Members.

**Muluk Member** is composed of dark grey clays with sand beds intercalation and restricted to the inner portion of Wadi El Natrun with exposed thickness of about 25 m at Quart El Muluk (type locality). However, the subsurface succession is composed of pyretic dark grey clays alternating with sand beds.

**El Solymanya Member** is composed of light green sand and sandstone alternating with shelly limestone and it is restricted to the periphery of Wadi El Natrun. It has 30 thick at Ras El Solymanya (type locality).

**El Hagif Formation** is mainly composed of thick white limestone beds intercalated with argillaceous sequence and it is restricted to the west of Wadi El Natrun depression. The thickness of El Hagif Formation ranges between 30 m and 40 m.

**Quaternary deposits:** Quaternary deposits cover wide stretches of the study area and it distinguished into different types as deltaic deposits and crust, alluvial deposits, lagoonal deposits, salt marshes and sabkha and Aeolian deposits. The thickness of the Quaternary sediments varies from 350 m to 25 m, it increases in the areas close to delta and decreases in the area close to Wadi El Natrun. Lithofacies of the Quaternary sediments are Sandy gravel, gravelly sand, slightly gravelly sand and sand. The lithofacies of the Quaternary sediments varies laterally and vertically and it has a direct impact on its hydraulic parameters. The environments of deposition are different all over the Quaternary sediments from turbidity currents, fluvial and beach. The fluvial environment is the dominant.

**Structure:** Folds, faults, unconformities and basaltic intrusions mainly affect the area of study. These types of structural elements are the most important factors controlling the groundwater conditions.

**Folds:** The area of study is mainly affected by two fold systems.

-The NE-SW folds' system (Syrian Arcs); the most conspicuous fold is Abu Roash domal structure.

-The NW-SE (clysmic) folds system; the most conspicuous folds are Wadi El Farigh and Wadi El Natrun anticlinal structure.

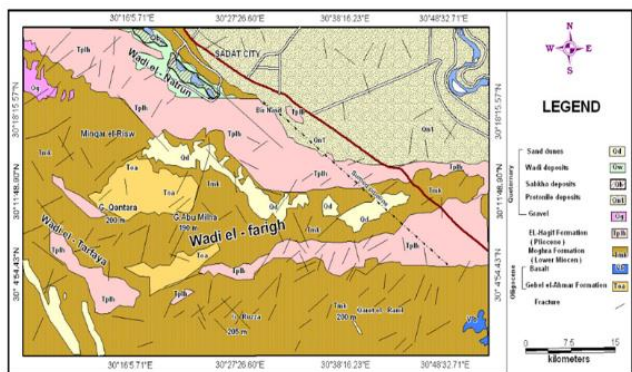
**Faults systems:** The study area is mainly affected by three normal faults systems. These systems in decreasing order of abundance are as follows:

- NW-SE (Clysmic) system

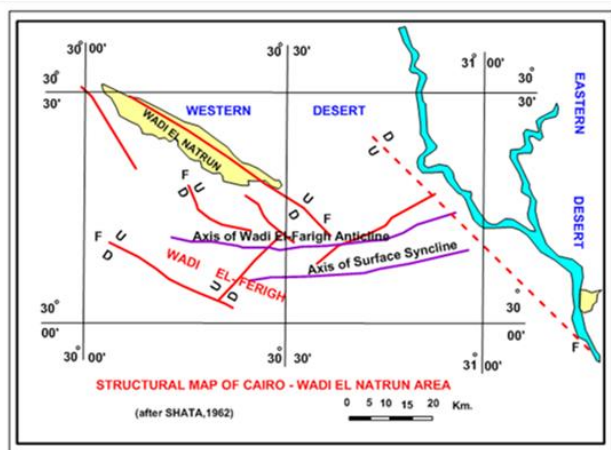
- NE-SW (Aqaba) system.

-E-W (Tethys) system.

The investigated area is separated from the Delta basin by one or more step faults having an eastern downthrown. Such faults are responsible for the facing of highly permeable Quaternary deposits of the Nile Delta with older sediments due west, and consequently, a westward flow of the Nile water to the western old sediments. Also, an important principle fault trending in a NW-SE, running approximately parallel to Rosetta branch is strongly expected to separate the Delta basin from the area to the west.



**Figure 3.** Geologic map of west Nile Delta, Egypt, (Conoco, 1987).



**Figure 4.** Structural map of Cairo-wadi El-Natron area after (Shata, 1962).

#### IV. The Hydrogeological Setting

From the hydrogeological point of view, special emphasis is given to the following topics:

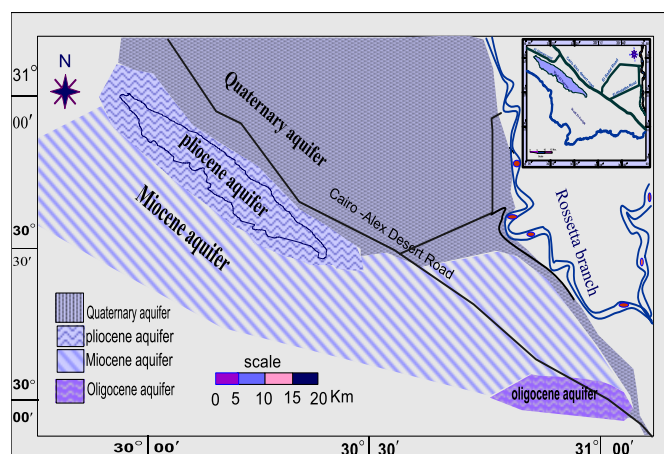
**Surface water system:** The surface water system is mainly Rosetta branch, El Rayah El Beheri, El Rayah El Nasserri and smaller canals. Such water system is subjected to infiltration to the groundwater aquifers. The canals are cutting the Nile silt and sandy clay deposits, discharging their water into the cultivated land.

**Groundwater aquifers:** The groundwater in the study area is mainly controlled by the geological conditions including lithology and geological structures. The main water bearing formations in the study area are:

- The Quaternary aquifer
- The Pliocene aquifer
- The Lower Miocene aquifer (Moghra), and
- The Oligocene aquifer

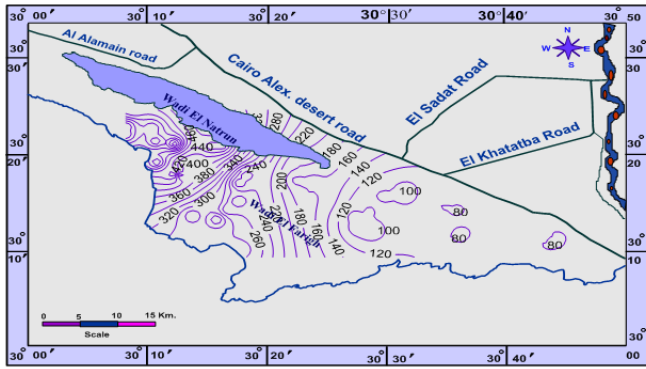
Miocene aquifer covers a wide area south and southwest of Wadi El Natrun and it is mainly composed of coarse sands and clay lenses intercalation (Fig. 5). The

saturated thickness of the Lower Miocene aquifer varies from 741 m in the western portion, 179 at K52 east Cairo-Alex. Desert Road and 16 m at Gebel Khashm El Kalb. Miocene aquifer has been deposited on the basaltic sheet at different levels; in the form of horst and graben-like structures. The amount of groundwater stored in Miocene aquifer is about  $25.9 \times 10^9 \text{ m}^3$  and the amount of withdrawal which equals the drawdown in groundwater is about  $0.036 \times 10^9 \text{ m}^3/\text{year}$ . This aquifer is under artesian conditions in the east and changes to water table aquifer towards the west. In Wadi El Natrun depression, where this aquifer is overlain by impervious Pliocene clay, the groundwater exists under confined condition. The ground water level in the Miocene aquifer varies between +8 m at the border with the Nile delta Quaternary aquifer to about -22.1 at the west. The high groundwater extraction allows the formation of the depression cones of piezometric levels. Isotopes indicate that the main recharging sources for the Miocene aquifer are old Nile water (before the construction of the High Dam) and small contribution of the recent Nile water in the area beside El Rayah El Nasserri. The estimated  $C^{14}$  age is 16737 YBP. The old age confirms the presence of ancient recharge component accompanying the recent one.



**Figure 5.** Distribution of different aquifers in the west Nile Delta area.

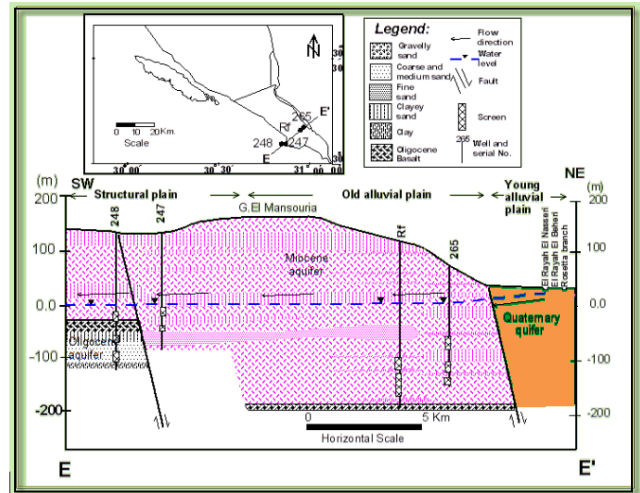
The total thickness of the Miocene sediments is varying from 75 m in the northeastern portion to about 250 m at Wadi - El Natrun, Regionally this thickness increases in northwest direction (El Ghazawi and Attwa, 1994). The saturated thickness of the Miocene aquifer (fig 6) is controlled by the prevailing structural conditions in the area. The maximum saturated thickness attains about 169m, and thins out in the northeastern direction to reach about 27 m at the Khashm El Kalb area.



**Figure 6.** Saturated thickness contour map of the Miocene aquifer after Shafeek (2015).

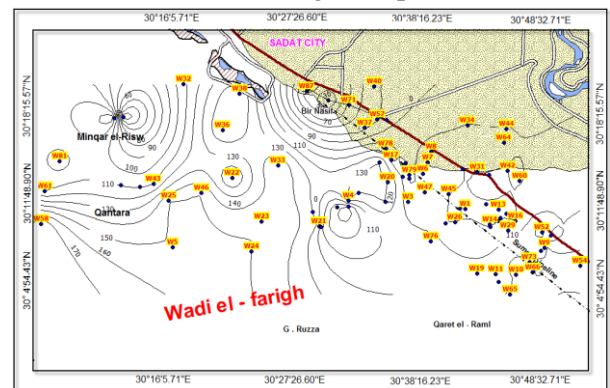
The southeastern boundaries of the lower Miocene aquifer are delimited by a set of faults with downthrows to the north where the Oligocene basaltic sheet is highly elevated (from + 150 m to + 30 m) above the regional potentiometric surface of the Lower Miocene aquifer (+2 m) (Fig. 7) accordingly, this aquifer is hydraulically connected with the underlying Oligocene aquifer. The northern boundary of Lower Miocene aquifer is bounded by many faults. South Wadi El Natron, this aquifer is bounded by NW- SE faults with their down throwing sides to the east where the Lower Miocene is uplifted in the front of the Pliocene aquifer.

The depth to water in Miocene aquifer varies from 49 m at the area close to El Rayah El Naseri to 176.88 m at high topographic areas towards the west (Fig. 8). The groundwater level in the Miocene aquifer varies between +6 m at the border with the Nile delta Quaternary aquifer to about -33.8 at the west. The ground water level at El Qattara depression is about -60 m below sea level accordingly, the groundwater flows to the west (El Qattara depression) with hydraulic gradient of about 20 cm/km (Hefny, et al, 1991). The groundwater in the Miocene aquifer flows westward, through Wadi El Farigh (Fig. 9). This movement is mainly controlled by an old buried Nile channels which divert groundwater flow towards Wadi El Natrun depression. There are two other groundwater flow directions one from the northeast (from Quaternary to Lower Miocene) and the second from southeast. High extraction rate along Cairo-Alexandria desert road and Wadi El Farigh allow the formation of the cones of depressions with high hydraulic gradient.

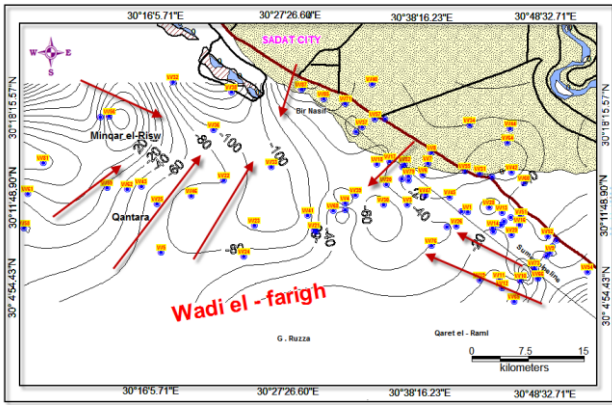


**Figure 7.** Hydrogeological Cross Section E-E' after Al Abd (2005).

The total amount of groundwater recharge from the Nile Delta aquifer is estimated to be in the rang of  $50$  to  $100 \times 10^6$  m<sup>3</sup>/year (RIGW/IWACO, 1990b). The amount of recharge was estimated using groundwater model as  $84 \times 10^6$  m<sup>3</sup>/year (Diab, et al. 1992). The discharge of this aquifer is mainly through the extraction from wells and is relatively discharged into the Pliocene aquifer at the southern part of Wadi El Natrun (Gomaa, 1995). High extraction rate ( $20$ - $30 \times 10^6$  m<sup>3</sup>/year/Km<sup>2</sup>) was found along Cairo-Alexandria desert road (El Fakhary, 1998). The extraction of large volumes of groundwater allow the formation of cones of the depressions of pizometric levels. Since the extraction rate are still increasing the lowering of water level and will continue in the future (RIGW/IWACO, 1990b). The environmental isotopes indicate that the main recharging source for the Miocene aquifer is the old Nile water before the construction of the High Dam with few contribution from recent Nile water in the area beside El Rayah El Naseri. The estimated age using C<sup>14</sup> is 16737 YBP (Dahab et al., 1998) confirming the dominance of ancient recharge component.



**Figure 4.** Depth to Water Contour Map of the Miocene Aquifer (2015).

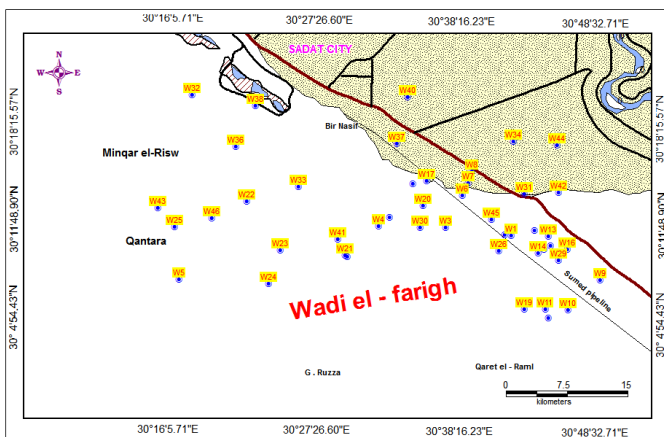


**Figure 9.** flow net map of the Miocene Aquifer

### V. Hydrochemical

The chemical composition of groundwater also reflects the ecosystem function, so that it is important to detect any changes resulting from natural systems and or caused by development processes. For the assessment of the chemical composition of the groundwater in the study areas 46 water samples from 46 water points were collected from wadi El-Farigh as showing as Fig. (10).

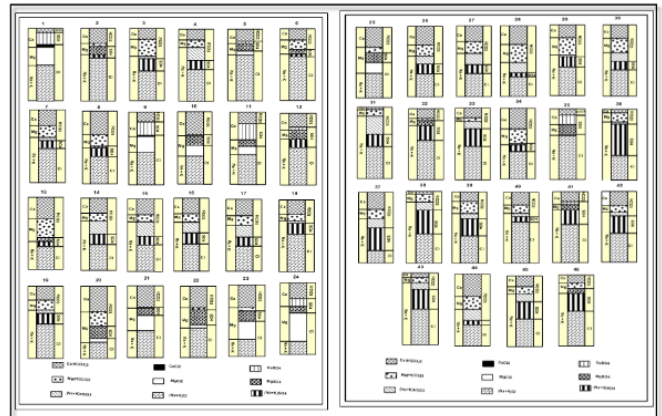
The total salinity of the Miocene aquifer (Moghra aquifer) varies from fresh to slightly brackish water. Local anomalies in water salinity are detected at the area East Cairo-Alex desert road, west El Rayiah El Nasserri and southwest Wadi El Natrun due to the over pumping rates, the presence of shallow clay lenses within the water bearing layers especially west El Rayiah El Nasserri and the low groundwater recharge. The total salinity ranges between 252 ppm and 2572 ppm. In Wadi El-Farigh the pH values vary from 7.11 to 8.5. All PH values of the selected water samples lie in the range of alkalinity.



**Figure 10.** Groundwater samples location map.

The hypothetical salt combinations of selected water samples are distinguished into six main assemblages as shown in Figures (11). Palmar (1911) as follows:

- **Group 1:** (Na+K) Cl, MgCl<sub>2</sub>, CaCl<sub>2</sub>, CaSO<sub>4</sub>, Ca (HCO<sub>3</sub>)<sub>2</sub>. (Samples No. w1 and w9).
- **Group 2:** (Na+K) Cl, (Na+K) SO<sub>4</sub>, MgSO<sub>4</sub>, Mg (HCO<sub>3</sub>)<sub>2</sub>, Ca (HCO<sub>3</sub>)<sub>2</sub>. (Samples No. w2, w6, w8, w13, w32, w34 and w46).
- **Group3:** (Na+K)Cl, (Na+K)SO<sub>4</sub>, (Na+K)HCO<sub>3</sub>, Mg(HCO<sub>3</sub>)<sub>2</sub>, Ca(HCO<sub>3</sub>)<sub>2</sub>. (Samples No. w3, w4, w7, w14, w15, w16, w17, w18, w19, w26, w27, w28, w29, w30, w31, w33, w36, w37, w38, w39, w40, w42, w43, w44, and w45).
- **Group4:** (Na+K) Cl, MgCl<sub>2</sub>, MgSO<sub>4</sub>, CaSO<sub>4</sub>, Ca (HCO<sub>3</sub>)<sub>2</sub>. (Samples No. w5, w10, w11, w21, w22, w24 and w35).
- **Group5:** (Na+K)Cl, (Na+K)SO<sub>4</sub>, MgSO<sub>4</sub>, CaSO<sub>4</sub>, Ca(HCO<sub>3</sub>)<sub>2</sub>. (Samples No. w12 and w41).
- **Group6:** (Na+K)Cl, MgCl<sub>2</sub>, MgSO<sub>4</sub>, Mg(HCO<sub>3</sub>)<sub>2</sub>, Ca(HCO<sub>3</sub>)<sub>2</sub>. (Samples No. w20, w23 and w25). w23 and w25).



**Figure 11.** Palmer diagram showing the hypothetical salt combination of the analyzed samples no. (1-46).

**Hydrochemical Classification Using Sunil's Diagram (1946)** show that four water types represent the water samples of Moghra Aquifer as shown in Figure (12). These are discussed hereinafter.

**A – The Na<sub>2</sub>SO<sub>4</sub> water type** is a mixed water of meteoric genesis it is represented by samples nos. (2, 6, 8, 12, 13, 29, 32, 35, 41 and 46) where (rNa/rCl) > 1 and (rNa/rK) - rCl/ rSo<sub>4</sub> < 1, which reflect the hydrochemical composition if water samples and the hypothetical salt combinations of this water and the hypothetical salt combinations of this water are (Na+K) Cl, (Na+K) SO<sub>4</sub>, MgSO<sub>4</sub>, Mg (HCO<sub>3</sub>)<sub>2</sub>, Ca (HCO<sub>3</sub>)<sub>2</sub>. These hypothetical salt combinations represent a chemical formation of infiltrating water of meteoric genesis.

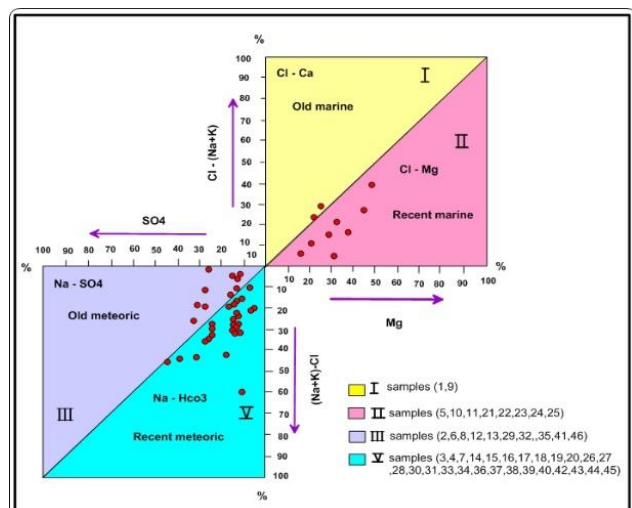
**B – The NaHCO<sub>3</sub> water type** water of meteoric water is represented by samples nos. (3, 4, 7, 14, 15, 16, 17, 18, 19, 20, 26, 27, 28, 30, 31, 33, 34, 36,

37,38,39,40,42,43,44 and 45). These samples are characterized by  $(rNa/rCl) > 1$  and  $(rNa/rK) - rCl / rSO_4 > 1$  which reflect a meteoric origin.

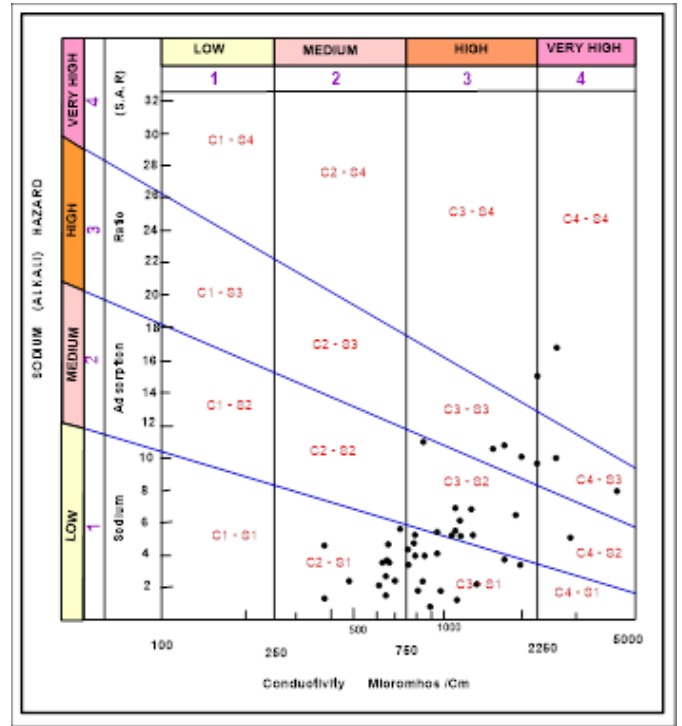
**C - The CaCl<sub>2</sub> water type is represented by two water samples nos. (1, 9) where  $(rNa/rCl) < 1$  and  $rCl - (rNa/rK) / rMg > .1$  This reflects an old marine origin. That is confirmed by the higher salinity of the two samples of about 1895 ppm.**

**D - The MgCl<sub>2</sub> water type is represented by rest of the water samples which are characterized by the ratio of  $(rNa/rCl) < 1$  and  $rCl - (rNa/rK) / rMg < 1$ . These water reflect a marine water origin. Their hypothetical salt combinations are (Na+K) Cl, MgCl<sub>2</sub>, MgSO<sub>4</sub>, Mg (HCO<sub>3</sub>)<sub>2</sub> and Ca (HCO<sub>3</sub>)<sub>2</sub>. These salts represent the chemical formation of normal composition of sea water.**

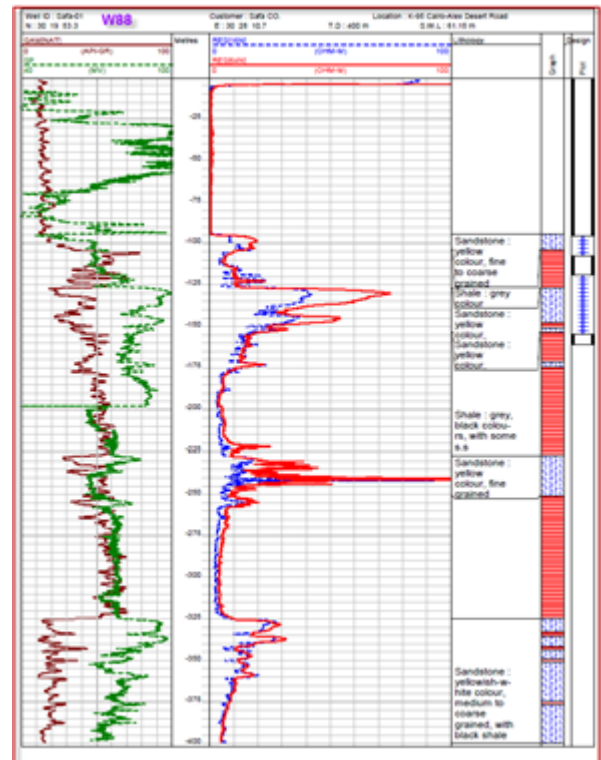
According to the U.S. Laboratory Salinity diagram (1954), the groundwater samples of Moghra aquifer system are of low sodium hazard and fall in the fields C<sub>2</sub>-S<sub>1</sub> and C<sub>3</sub>-S<sub>1</sub>, where S<sub>1</sub> is low sodium water, C<sub>2</sub> is moderate salinity water which is good for soils of medium permeability for most plants, and other sample lies in the C<sub>3</sub>-S<sub>2</sub>, C<sub>4</sub>-S<sub>2</sub>, C<sub>3</sub>-S<sub>3</sub>, C<sub>4</sub>-S<sub>3</sub> and C<sub>4</sub>-S<sub>4</sub> field. From Classified as groundwater in terms of sodium absorption ratio method U.S. Laboratory Salinity diagram (1954), it clear that the samples lie near the medium salinity water, which is satisfactory for plants having moderate salt tolerance on soils of moderate permeability with leaching. The samples fall in fields C<sub>3</sub>-S<sub>3</sub>, C<sub>4</sub>-S<sub>3</sub> and C<sub>4</sub>-S<sub>4</sub> high to vary high sodium hazard and salinity water. Figures (13),

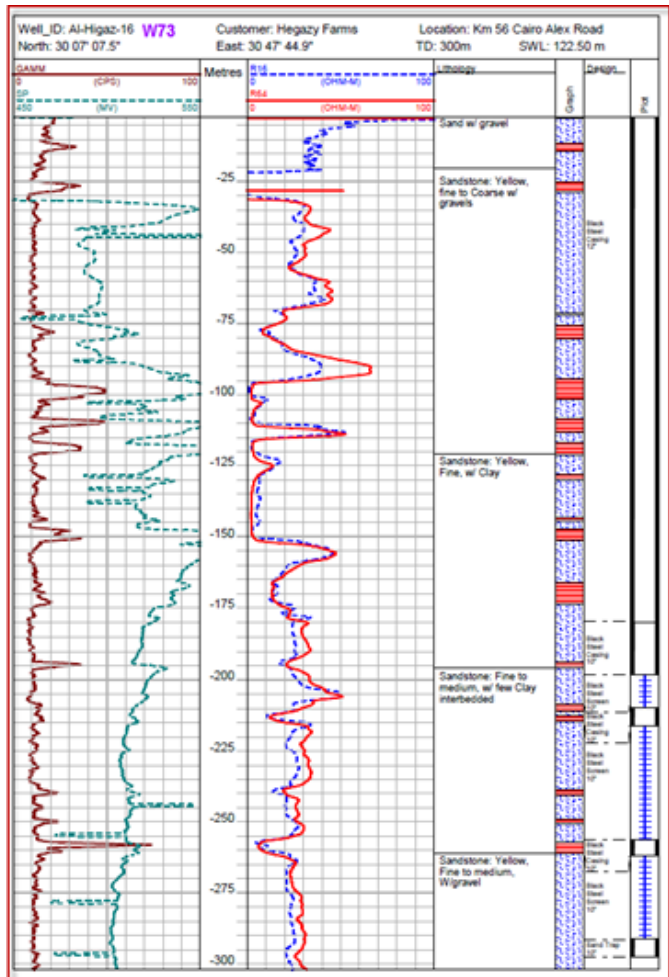


**Figure 12.** Sulin's graph for genetic classification of the groundwater samples of the Moghra aquifer.



**Figure (13):** Diagram for use in interpreting the analysis of irrigation water. Adapted from U.S. Laboratory Salinity staff (1954).





Well No.	T.D.S ppm	E.C. micromhos/cm	PH	Units	Cations					Sum.	Anions			S.A.R	Ht. Epm	R.S.C Epm
					Ca	Mg	Na+K	Cl	SO4		CO3	NO3	Sum.			
W1	1895	2960	8.5	Epm	39.1	22.99	20.04	12.16	61.02	35.45	48.03	20.24034	11.821	713.575	-12.8521	
W2	779	1217	8.41	Epm	4.5	180	48	23	175	344	85	11.520517	5.31714	216.8	-1.48804	
W3	618	965	8	Epm	6.6	434	46.64	19	267	133	24	9.93561	190.075	10.01612		
W4	463	724	7.94	Epm	2.3	123.6	23.33	8.42	190	114.8	48	7.269772	91.947	1.274691		
W5	1209	1888	7.8	Epm	0.05824	5.371991	1.064274	0.774671	3.277644	13.20169	0.999275	7.351472	6.562052	-3.34319		
W6	721	1126	7.94	Epm	0.08957	1.734519	1.236227	1.138158	3.051793	6.792666	1.207519	11.03204	3.125096	-0.84289		
W7	538	840	8.01	Epm	0.08954	1.680513	1.048018	1.334704	3.665375	4.096642	0.905864	5.516701	4.11131	0.429853		
W8	326	509	8.06	Epm	0.09429	2.386246	1.664074	0.919184	2.474808	1.941896	0.664148	5.006061	2.085039	-0.10136		
W9	1174	1836	8.1	Epm	0.26397	0.846877	0.237326	1.419442	2.228777	12.12976	3.788478	10.12702	3.537464	-1.42069		
W10	536	837	7.94	Epm	0.12532	3.048822	2.041338	1.247105	3.566539	4.372555	1.29688	8.219754	1.897774	-2.5817		

Well No.	T.D.S ppm	E.C. micromhos/cm	PH	Units	Cations					Sum.	Anions			S.A.R	Ht. Epm	R.S.C Epm
					Ca	Mg	Na+K	Cl	SO4		CO3	NO3	Sum.			
W11	1299	2030	7.9	Epm	0.19001	0.532971	7.515097	1.118142	3.113732	10.38802	0.100254	20.00444	19.59491	3.353889	-6.44681	
W12	893	1391	7.97	Epm	0.107417	0.654953	3.113273	1.606886	2.638479	6.911422	3.538455	13.39132	10.08908	5.575454	-2.00888	
W13	420	657	8.01	Epm	0.089544	2.432042	1.996088	1.973684	3.654539	2.11566	0.845305	48.79093	37.21967	15.96229	6.15515	
W14	506	790	7.94	Epm	0.179028	5.393451	1.806239	0.881579	3.900361	2.885750	1.148954	20.72062	36.40386	14.39353	1.318143	
W15	433	677	8	Epm	0.153453	4.784689	1.48964	0.705992	3.896659	2.578729	0.805746	1.847123	12.51966	14.77841	4.576992	
W16	492	769	7.97	Epm	0.240684	5.132666	1.420216	0.934219	3.834888	2.502247	1.099373	20.15787	32.80084	4.437073	1.156892	
W17	428	668	8.06	Epm	0.153453	4.784689	1.48964	0.705992	3.896659	2.578729	0.805746	1.847123	12.51966	14.77841	4.576992	
W18	816	1214	7.97	Epm	0.148016	0.671926	1.517350	1.0331	3.379664	5.37258	12.81878	13.08657	12.73371	6.32863	0.378248	
W19	763	1192	7.98	Epm	0.130205	7.650462	2.280938	1.842237	4.319682	5.531373	1.557136	11.63388	11.39916	5.908206	0.446897	
W20	292	399	8.17	Epm	0.132992	1.953773	3.143173	1.536184	2.815504	1.156559	0.748711	6.710202	11.59629	1.617055	-4.41834	

Well No.	T.D.S ppm	E.C. micromhos/cm	PH	Units	Cations					Sum.	Anions			S.A.R	Ht. Epm	R.S.C Epm			
					Ca	Mg	Na+K	Cl	SO4		CO3	NO3	Sum.						
W21	876.8	1370	7.76	Epm	0.2	150	45	14.1	11.72	114.0002	80.7004	83.6296	242.0508	207.2514	78.548	13.7	2.427163	-4.44	
W22	969.6	890	7.9	Epm	2.82	75.887	46.2044	31.4988	2.82	75.887	46.2044	31.4988	182.0702	141.1178	14.3222	0.4	2.108293	-1.89	
W23	704	1100	7.69	Epm	0.9	2.36	3.52	4.97	0.9	2.36	3.52	4.97	26.5448	108.2395	42.7481	11	1.14815	-4.45	
W24	588.5	920	7.12	Epm	0.24	32.408	59.7192	44.3019	0.24	32.408	59.7192	44.3019	32.6087	58.71729	19.66566	9.2	0.793797	-4.43	
W25	472	1059	7.87	Epm	0.33	3.35	3.9	3.9	0.33	3.35	3.9	3.9	25.1424	108.594	62.9193	10.5	1.748741	-3.03	
W26	600	650	7.8	Epm	0.333333	30.95238	28.51413	37.14286	0.333333	30.95238	28.51413	37.14286	36.86744	10.66667	12.47619	11	1.54855	-4.45	
W27	500	665	7.4	Epm	0.148016	0.671926	1.517350	1.0331	0.148016	0.671926	1.517350	1.0331	0.148016	0.671926	1.517350	1.0331	0.148016	0.671926	1.517350
W28	630	700	7.7	Epm	0.179028	5.393451	1.806239	0.881579	0.179028	5.393451	1.806239	0.881579	0.179028	5.393451	1.806239	0.881579	0.179028	5.393451	1.806239
W29	600	780	7.7	Epm	0.148016	0.671926	1.517350	1.0331	0.148016	0.671926	1.517350	1.0331	0.148016	0.671926	1.517350	1.0331	0.148016	0.671926	1.517350
W30	620	800	7.3	Epm	0.063708	0.567838	1.711717	1.024342	0.063708	0.567838	1.711717	1.024342	0.063708	0.567838	1.711717	1.024342	0.063708	0.567838	1.711717
W31	650	850	7.5	Epm	0.051151	0.4689	0.199661	0.096236	0.051151	0.4689	0.199661	0.096236	0.051151	0.4689	0.199661	0.096236	0.051151	0.4689	0.199661

Well No.	T.D.S ppm	E.C. micromhos/cm	PH	Units	Cations					Sum.	Anions			S.A.R	Ht. Epm	R.S.C Epm				
					Ca	Mg	Na+K	Cl	SO4		CO3	NO3	Sum.							
W32	1334	2610	7.78	Epm	0.12302	1.42	80	20	0.12302	1.42	80	20	214	119	288	21.32276	16.21031	264.8	-1.7396	
W33	1058	1590	7.9	Epm	0.154553	13.68152	2.936609	6.749132	0.154553	13.68152	2.936609	6.749132	2.827825	46.30278	24.82047	15.8379	10.77026	146.9	1.570979	
W34	421	610	7.9	Epm	0.179028	2.781013	1.392206	1.609211	0.179028	2.781013	1.392206	1.609211	1.95474	2.59579	0.490868	6.20584	2.267205	160.2	-0.1074	
W35	2572	4000	7.53	Epm	0.154553	26.01131	9.191427	6.332237	0.154553	26.01131	9.191427	6.332237	9.198818	62.94471	27.80642	41.31533	3.333539	775.7	-11.7418	
W36	1158	1700	7.81	Epm	0.220179	13.70161	0.796403	2.481105	0.220179	13.70161	0.796403	2.481105	3.507407	6.094453	7.516136	10.07163	10.72287	163	0.241538	
W37	515	420	8.15	Epm	0.785242	42.47319	22.8896	12.83841	0.785242	42.47319	22.8896	12.83841	49.67988	19.42577	20.63025	6.03793	4.518889	174.5	-0.52995	
W38	1520	2250	8.07	Epm	0.220179	19.22574	0.980004	2.220285	0.220179	19.22574	0.980004	2.220285	25.01111	41.54096	33.44791	22.40803	15.15578	2.238321		
W39	354	1320	8.1	Epm	0.127877	0.646372	1.846307	2.302622	0.127877	0.646372	1.846307	2.302622	3.014749	5.507025	3.252072	28.14182	39.66608	24.1721	0.8581	
W40	567	800	7.91	Epm	0.179028	0.671926	1.517350	1.0331	0.179028	0.671926	1.517350	1.0331	2.827825	5.01157	0.603789	4.27206	5.199174	0.136475		
W41	1744	2630	7.11	Epm	0.858	73.8414	16.14037	17.67191	0.858	73.8414	16.14037	17.67191	13.07186	56.18419	26.80654	24.31	13.8733	6.704409	307.3	0.8581
W42	648	950	8.14	Epm	0.240684	0.513628	1.396208	0.888482	0.240684	0.513628	1.396208	0.888482	3.130721	3.892097	2.208254	19.1	3.228832	5.651641	0.247072	
W43	1626	2430	8.01	Epm	0.220179	20.31318	4.93002	1.973844	0.220179	20.31318	4.93002	1.973844	2.827825	12.51966	14.77841	11.32851	24.24853	16.95077	2.089252	
W44	487	660	8.05	Epm	0.102022	3.249262	2.095899	1.551316	0.102022	3.249262	2.095899	1.551316	4.93002	1.973844	11.32851	6.50544	2.62056	0.948222		
W45	799	1170	8.19	Epm	0.220179	0.525446	1.746202	1.315789	0.220179	0.525446	1.746202	1.315789	2.827825	4.437073	2.210154	11.40564	6.88983	0.969169		
W46	1462	2220	7.79	Epm	0.154553	15.22962	3.225261	2.52929	0.154553	15.22962	3.225261	2.52929	3.507407	6.094453	7.516136	22.25298	19.24688	277.6	-1.45685	

## VI. Conclusion

The ground water in wadi El- Farigh is available from an aquifer system belonging to Lower Miocene times (Moghra aquifer.). The total thickness of the Miocene sediments is varying from 75 m in the northeastern portion to about 250 m at Wadi - El Natron, Regionally this thickness increases in northwest direction. The depth to water in Miocene aquifer varies from 49 m at the area close to El Rayah El Naseri to 176.88 m at high topographic areas towards the west. The groundwater in the Miocene aquifer flows westward, through Wadi El Farigh. This movement is mainly controlled by an old buried Nile channels which divert groundwater flow towards Wadi El Natron depression.

The total salinity of the Miocene aquifer (Moghra aquifer) varies from fresh to slightly brackish water. The total salinity ranges between 252 ppm and 2572 ppm. In Wadi El-Farigh the pH values varies from 7.11 to 8.5. All PH values of the selected water samples lie in the range of alkalinity.

According to the U.S. Laboratory Salinity diagram, the groundwater samples of Moghra aquifer system are of low sodium hazard and fall in the fields C<sub>2</sub>-S<sub>1</sub> and C<sub>3</sub>-S<sub>1</sub>, where S<sub>1</sub> is low sodium water, C<sub>2</sub> is moderate salinity



water which is good for soils of medium permeability for most plants, and other sample lies in the C<sub>3</sub>-S<sub>2</sub>, C<sub>4</sub>-S<sub>2</sub>, C<sub>3</sub>-S<sub>3</sub>, C<sub>4</sub>-S<sub>3</sub> and C<sub>4</sub>-S<sub>4</sub> field. From Classified as groundwater in terms of sodium absorption ratio U.S. Laboratory Salinity diagram clear that the samples lie near the medium salinity water, which is satisfactory for plants having moderate salt tolerance on soils of moderate permeability with leaching. The samples fall in fields C<sub>3</sub>-S<sub>3</sub>, C<sub>4</sub>-S<sub>3</sub> and C<sub>4</sub>-S<sub>4</sub> high to vary high sodium hazard and salinity water.

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